





Geotechnical Module in REDAS

Landslide Hazard

Liquefaction

Nikolaos Klimis¹,

K. Papatheodorou², S. Valkaniotis¹, El. Petala¹, G. Papathanassiou³ and J. Gkiougkis¹

¹ Civil Engineering Department, Democritus University of Thrace (DUTh)

² Survey & Geoinformatics Engineering Department, International Hellenic University (IHU)

³ Geology Department, Aristotle University of Thessaloniki (AUTh)



In the event of a strong earthquake part of damages are related to

ground failures

We focus on two of them (at a regional scale):

Landslides
Liquefaction

The geotechnical module is part of the REDA system and it has been tested at pilot regions where earthquake triggered landslides and liquefaction phenomena occurred.



In the present project **2 different methodological approaches for LHA** have been selected:

- ➤ The 1st belongs to the physically based methods,
- ➤ the 2nd one belongs to the statistically based methods

Cons and Pros:

- Statistical methods are oriented to regional scale where spatial variability of geological formations and engineering parameters are important and landslide type relatively complex
- Physically based methods are more engineer oriented and better conceived in engineered projects and decision making, provided that geological and geomorphological conditions are fairly homogeneous and landslide types relatively simple, even if inventories are incomplete or inexistant Common borders. Common solutions.



Nowicki et al. (2014) developed an empirical landslide probability model, combining **shaking estimates** with landslide susceptibility proxies, i.e., **topographic slope, surface geology and climate parameters**.

Jessee et al. (2018) proposed an updated model regarding the near-real time assessment of seismically induced landslides resulting in the following best fitting model

 $t = \text{Logit}(P) = a + b \times \ln(\text{PGV}) + c \times Slope + d \times Lithology$ $+ e \times Land Cover + f \times CTI + g \times \ln(\text{PGV}) \times Slope,$

the values of coefficients: a, b, c, d, e, f and g are solved within the regression

The **predicted probability of landslide occurrence P(t)** can be computed based on the following formula:

P(t)=1/(1+exp(-t))





Scenario:

The November 17th, 2015 Lefkada (Greece) strike-slip Mw 6.4 earthquake

Based on the updated best fitting statistical model of Jessee et al. (2018):

- Peak Ground Velocity (PGV)
- CTI (soil wetness)

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- Lithology
- Land cover
- Slope (DEM)



Papathanassiou et al. 2017



Scenario: 2015 Lefkada

Input layers for statistical landslide model (Jessee et al. 2018)

- Peak Ground Velocity (PGV)
- Vs30
- Lithology
- Slope (DEM)
- CTI (soil wetness)

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Land cover





CROSS BORDER

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X

Landslide Hazard (statistical model)





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Statistical landslide probability model (Jessee et al. 2018)

Proposed additional constraints and thresholds:

Slope > 5 degrees

Spatial Resolution:

 High resolution (e.g. 0.00025 deg) offers better spatial resolution and takes advantage of available high resolution Digital Elevation Models





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Statistical landslide probability model (Jessee et al. 2018)





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Statistical landslide probability model (Jessee et al. 2018)





20.50

38.80

38.70

+ 38.60

Probability

0 - 0.2

0.2 - 0.4

0.4 - 0.6

0.6 - 0.8

0.8 - 1

Mw 6.4

 \checkmark

20151117

Landslide Hazard (statistical model)

Landslide Statistical Model Output

P(t)=1/(1+exp(-t))

20.60

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20.70



 $L_{P}(P) = e^{(a+b\times P+c\times P^{2}+d\times P^{3})},$ a = -7.592 b = 5.237, c = -3.042,

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d = 4.035

This equation corrects the predicted probability. Represents the **frequency of landslide occurrence** i.e. the portion of each cell expected to have landslide occurrence (areal coverage)



Landslide Statistical Model Output – Validation with surveyed co-seismic landslides





Landslide Statistical Model Output – Validation with surveyed co-seismic landslides





Physically based landslide hazard assessment methods are based on the modelling of **slope failure processes and the output is the F_S values**.

Use of the **infinite slope model**, is suitable for **shallow earthquake induced landslides**. The driving equation is:

$$F = \frac{c' + (z\gamma\cos^2\beta - z\rho\omega\cos\beta\sin\beta - \gamma_w z_w\cos^2\beta)\tan\phi'}{z\gamma\sin\beta\cos\beta + z\rho\alpha\cos^2\beta}$$

φ': effective angle of friction of geomaterial (⁰) c': effective cohesion of geomaterial (kPa), β: slope angle (Deg), ρ: bulk density (Kg/m³) γ: specific weight (kN/m³), $γ_w$: specific weight of the water (kN/m³), *a*: earthquake acceleration (m/s²) *z*: normal thickness of the failure slab (m) $m = z_w/z$ % of the water saturated failure slab





Landslide Hazard (Infinite Slope Model)







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Landslide Hazard

Wet conditions Sliding mass thickness: 1m



Statistical model and observed landslides





Landslide Hazard (Infinite Slope Model)

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Blind test prediction northern of Serres mountainous area

Sliding mass thickness: Variable based on slope and lithology (0-2m)

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Probability of liquefaction based on Zhu et al (2017) methodology

Probability of liquefaction

$$P(x) = \frac{1}{1 + e^{-x}}$$

Coastal region (dc<20km)

Coastal

42.08

62.59

11.43

 $x = 12.345 + 0.301 \times \ln(PGV) - 2.615 \times \ln(Vs_{30}) + 5.556 \times 10^{-4} \times Precip - 0.0287 \times \sqrt{dc} + 0.0666 \times dr - 0.0369 \times (\sqrt{dc} \times dr)$

Non Coastal region (dc>20km)

 $x = 8.801 + 0.334 \times \ln(PGV) - 1.918 \times \ln(Vs_{30}) + 5.408 \times 10^{-4} \times Precip - 0.2054 \times dw \\ - 0.0333 \times wtd$

Spatial extent of liquefaction

42.4

9.165

Where:

PGV (cm/s): Peak ground velocity VS₃₀ (m/s): Shear wave velocity to 30m depth Precip (mm): Precipitation (mean annual) dc (km): Distance to the nearest coast dr (km): Distance to the nearest river dw (km): Distance to the nearest water body wtd (m): Water table depth

$$SF = \frac{1}{1 + exp\left(-2[M-6]\right)}$$

for M (earthquake magnitude) we use Mw

Common borders. Common solutions.



Where

Parameters

а

b

С



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M6.3 Larisa, Greece (03/03/2021)

Input Data

Vs₃₀ (m/s)

Precipitation (mm)

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Input Data

Dw (km)

Wtd (m)



Common borders. Common solutions.





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Probability of liquefaction - P(x)

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Probability of liquefaction - P(x)



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